

# [Turkish Journal of Earth Sciences](https://journals.tubitak.gov.tr/earth)

[Volume 33](https://journals.tubitak.gov.tr/earth/vol33) [Number 1](https://journals.tubitak.gov.tr/earth/vol33/iss1) SI-1

[Article 5](https://journals.tubitak.gov.tr/earth/vol33/iss1/5) 

1-4-2024

# Investigation of earthquake sequence and stress transfer in the Eastern Anatolia Fault Zone by Coulomb stress analysis

HAMDİ ALKAN hamdialkan@yyu.edu.tr

AYDIN BÜYÜKSARAÇ absarac@comu.edu.tr

ÖZCAN BEKTAŞ obektas@cumhuriyet.edu.tr

Follow this and additional works at: [https://journals.tubitak.gov.tr/earth](https://journals.tubitak.gov.tr/earth?utm_source=journals.tubitak.gov.tr%2Fearth%2Fvol33%2Fiss1%2F5&utm_medium=PDF&utm_campaign=PDFCoverPages) 

**C** Part of the Earth Sciences Commons

# Recommended Citation

ALKAN, HAMDİ; BÜYÜKSARAÇ, AYDIN; and BEKTAŞ, ÖZCAN (2024) "Investigation of earthquake sequence and stress transfer in the Eastern Anatolia Fault Zone by Coulomb stress analysis," Turkish Journal of Earth Sciences: Vol. 33: No. 1, Article 5.<https://doi.org/10.55730/1300-0985.1898> Available at: [https://journals.tubitak.gov.tr/earth/vol33/iss1/5](https://journals.tubitak.gov.tr/earth/vol33/iss1/5?utm_source=journals.tubitak.gov.tr%2Fearth%2Fvol33%2Fiss1%2F5&utm_medium=PDF&utm_campaign=PDFCoverPages) 

This Article is brought to you for free and open access by TÜBİTAK Academic Journals. It has been accepted for inclusion in Turkish Journal of Earth Sciences by an authorized editor of TÜBİTAK Academic Journals. For more information, please contact [academic.publications@tubitak.gov.tr](mailto:academic.publications@tubitak.gov.tr).



**Turkish Journal of Earth Sciences** Turkish J Earth Sci

**http://journals.tubitak.gov.tr/earth/**

**Research Article**

(2024) 33: 56-68 © TÜBİTAK doi:10.55730/1300-0985.1898

# **Investigation of earthquake sequence and stress transfer in the Eastern Anatolia Fault Zone by Coulomb stress analysis**

**Hamdi ALKAN1,\*, Aydın BÜYÜKSARAÇ2 , Özcan BEKTAŞ<sup>3</sup>**

<sup>1</sup> Department of Geophysics, Engineering Faculty, Van Yüzüncü Yıl University, Van, Turkiye <sup>2</sup>Can Vocational School, Çanakkale Onsekiz Mart University, Çanakkale, Turkiye  $3$  Department of Geophysics, Engineering Faculty, Sivas Cumhuriyet University, Sivas, Turkiye



**Abstract:** A devastating earthquake with a magnitude of  $(M_w = 7.7)$  occurred on February 06, 2023, in the Pazarcık segment of the Eastern Anatolian Fault Zone, which has not shown major earthquake activity for a long time. On the same day, another earthquake with a magnitude of ( $M_w$  = 7.6) occurred in Ekinözü-Elbistan (Kahramanmaraş) in the northwest. Three more earthquakes with magnitudes of  $M_{_{\rm w}}$  = 6.6,  $M_{_{\rm w}}$  = 5.9, and  $\rm M_{_{\rm L}}$  = 5.7 occurred on the same day, and significant damage, loss of life, and property occurred in 11 provinces and districts. A sixth earthquake occurred with  $M_w = 6.4$  magnitude in Defne-Hatay on February 20, 2023. In addition, more than 32,000 aftershocks were recorded while this study was in progress. The interconnection of these earthquakes was investigated by Coulomb stress analysis. It showed compatibility with the distribution of both these six earthquakes and small aftershocks, in which earthquakes transfer stress to each other during their formation phases. By calculating Coulomb stress changes, mainshock ruptures have played an important role in transferring stresses between Elazığ and Malatya in the northeast, Kahramanmaraş-Göksun in the west, and Hatay and Syria in the south. Also, our calculations showed that shallow depths in and around source zones received positive stress changes (~1.0 bar) due to the focal depths of mainshocks with their aftershocks. Finally, it has been better understood with these earthquakes that the stress transfer time of high-energy earthquakes can be shortened considerably.

**Keywords:** South Anatolia, coulomb stress analysis, earthquake, stress transfer

#### **1. Introduction**

On February 06, 2023, at 01:17:32 (UTC) in the southeast of Turkey, an earthquake with a magnitude of  $M_{\nu} = 7.7$ occurred, which is the second-largest earthquake in Turkey in the instrumental period. The magnitude of this earthquake was measured as  $M_w = 7.8$  by the USGS. In fact, this earthquake marked the beginning of a series of large earthquakes that occurred one after the other. Because right after, there was another aftershock with a magnitude of  $M_w = 6.6$  at 01:28:16 (UTC). Then, on the same day approximately nine h later from the first main shock at 10:24:47 (UTC), a new mainshock of  $M_v = 7.6$ , the fourth shock of  $M_{\rm L}$  = 5.7 at 10:51:28 (UTC), the fifth shock of  $M_w = 5.9$  at 12:02:11 (UTC), and the sixth mainshock of *M<sub>w</sub>* = 6.4 at 17:04:28 (UTC) occurred on February 20, 2023 (AFAD, 2023). As a result of these earthquakes, more than 50,000 people died in 11 provinces, more than 100,000 people were injured and more than 500,000 houses became uninhabitable. Considering the fault systems in which these earthquakes occurred, the fact that

almost all of them are on a different fault segment raises the possibility of different mainshocks, and since this region is the East Anatolian Fault Zone (EAFZ), it can be considered as an earthquake sequence triggered by a single movement (Figure 1). The place where the first earthquake  $(M_{\nu} = 7.7)$  occurred is located within the EAFZ, as well as at the northern end of the Dead Sea Fault Zone (DSFZ) (Figure 2). As a result of the complex tectonic structure of the region, earthquakes have occurred in relation to each other. Due to its left-lateral strike-slip and the approximately 580 km long EAFZ character, it transmits stress to the next segment after each earthquake that occurs in its segments. The tectonic system that Turkey is in has led to the formation of the EAFZ and the North Anatolian Fault Zone (NAFZ), and they have been the main reason for the westward movement of the Anatolian Plate geodynamically (Şengör et al. 1985; McClusky et al. 2000; Bozkurt 2001; Reilinger and McClusky, 2011).

Many researchers have conducted studies at different levels, mostly drawing attention to earthquake hazards



<sup>\*</sup> Correspondence: hamdialkan@yyu.edu.tr



Figure 1. Equal distributions of earthquakes data of  $\frac{1}{2}$ Figure 1. Epicentral distributions of earthquakes dated the 06<sup>th</sup> of February 2023 and 23<sup>rd</sup> of February 2023.

in the EAFZ. For example, Bayrak et al. (2015) calculated the *b*-value along the EAFZ and determined that the lowest *b*-value was around Karlıova. Irmak et al. (2021) concluded that the westward movement of the Central Anatolian High Plateau of January 24, 2020, Sivrice (Elazığ) earthquake ( $M_w = 6.8$ ) included not only the left lateral strike-slip movement of the EAF but also the extensional transtensional movements of the upper crust within and along the EAFZ. The expectation of an important earthquake in the EAFZ has been revealed in many studies. Lower *b*-values were observed in the NE section of the EAFZ compared to the SW section. Although the relatively high *b*-value in the Gölbaşı-Türkoğlu segment, which is considered to be a seismic gap in particular, creates a contradiction, the reason for this is that there were many low-magnitude earthquakes in the region between the dates studied (Kartal and Kadirioğlu, 2013). Gülerce et al. (2017) created the probabilistic seismichazard assessment maps for EAFZ. They calculated *b*-values around the rupture systems along the EAFZ. The low *b*-values (~0.81) were generally obtained for the whole of EAFZ, the smaller of which  $(-0.6)$  was calculated for the Ilıca-Karlıova region which has higher stress levels. Considering the stress transfer and future hazard forecast, Öztürk (2020) inferred that the seismic *b*-values showed a decreasing trend  $(-0.55)$  calculated in and around the Central Anatolian region, especially Kahramanmaraş. On the other hand, significant seismic quiescence areas included the west of Malatya and Kahramanmaraş. Coulomb analysis conducted by Alkan et al. (2021) after

the January 24, 2020, Sivrice (Elazığ) earthquake (*M*<sup>w</sup> = 6.8) related to stress transfer and segments triggering each other in the EAFZ, stated that the next earthquakes are expected between Adıyaman, Diyarbakır and Elazığ, Bingöl in the northeast of Kahramanmaraş. In the study carried out by examining the focal depths of earthquakes occurring in the EAFZ, it was observed that small and medium-sized earthquakes migrated from the main fault to adjacent fault segments systematically (Bulut et al., 2012). Duman and Emre (2013) determined that the Pötürge, Pazarcık, and Amanos fault segments along the EAFZ have the potential to produce large earthquakes in the near future. In this study, the Coulomb stress analysis method was used to determine the expected earthquake hazard locations in the EAFZ. The results of the method are sensitive to the history of major earthquakes in the region, and detailed studies are needed to define the exact rupture geometries of previous earthquakes in these segments (Nalbant et al., 2002).

#### **2. Seismotectonics of the East Anatolian Fault Zone**

When the tectonic system in Turkey is examined, the Arabian Plate, which is a part of the African Plate, moves north to the Eurasian Plate, compressing Eastern Anatolia, and with the effect of this compression, the NAFZ and the EAFZ effect, it is seen that the Anatolian Plate moves westward (Figure 2). The EAFZ, with an average length of 580 km, is one of the most seismically active regions of Turkey and many major earthquakes have occurred along this fault zone. The EAFZ is a left-lateral strike-slip fault, forming



Sea Basin, *MBSR*: Middle Black Sea Ridge, *WBSB*: West Black Sea Basin, *NEAFZ*: Northeast Anatolian Fault Zone, *WAGS*:<br>West Anatolian Graben System, *KTJ*: Karlıova Triple Junction, *EAFZ*: East Anatolian Fault Zone, *B* with the East Anatolian Fault Zone (modified from Reilinger et al. (2006) and Emre et al. (2018)). The black lines delineate the<br>major active faults. The dark red circles denote city centers. Abbreviations; AF.S: Afrin Seg *ER.*S: Erkenek Segment, GS: Gerger Segment, *H.FZ:* Harran Fault Zone, *KBFZ:* Kavakbaşı Fault Zone, *KM.FZ: K*ahramanmaraş<br>Fault Zone, *KD.F: K*aracadağ Fault, *KF: Ka*rataş Fault, *KO.S: K*ozluk Segment, *KS: Karl*ıova Savur Fault, SB.F: Saimbeyli Fault, *SS*: Sermada Segment, *SU.F*: Sürgü Fault, *YF*: Yumurtalık Fault, *YS*: Yesemek Segment. **Figure 2.** (a) Active tectonic map of Arabian and Anatolian Plates (modified from Okay and Tüysüz (1999), Alkan et al. (2021)). Yellow arrows indicate plate velocities (mm/yr). The red rectangle shows the study region. Abbreviations; *EBSB*: East Black Sea Basin, *MBSR*: Middle Black Sea Ridge, *WBSB*: West Black Sea Basin, *NEAFZ*: Northeast Anatolian Fault Zone, *WAGS*: Zone, *DSFZ*: Dead Sea Fault Zone, *EFZ*: Ecemiş Fault Zone. (b) The main tectonic elements and active fault segments along with the East Anatolian Fault Zone (modified from Reilinger et al. (2006) and Emre et al. (2018)). The black lines delineate the Bozova Fault, *CF*: Çardak Fault, *CU.S*: Çürgüş Segment, *IS*: Ilıca Segment, *L-D.S*: Lice-Dicle Segment, *E.FZ*: Engizek Fault Zone, *ER.S*: Erkenek Segment, *GS*: Gerger Segment, *H.FZ*: Harran Fault Zone, *KBFZ*: Kavakbaşı Fault Zone, *KM.FZ*: Kahramanmaraş Narince Segment, *NAFZ*: North Anatolian Fault Zone, *PA.S*: Palu Segment, *PAZ.S*: Pazarcık Fault, *PU.S*: Pütürge Segment, *SA.F*:

the southeast boundary of the Anatolian plate and intersecting the NAFZ at Karlıova. Similarly, the EAFZ merges with the DSFZ around Antakya (Figure 2). On the other hand, it basically extends from Karlıova to İskenderun Bay in the southwest direction with seven different segments (Figure 3) (Emre et al., 2018; Alkan et al., 2021). Large and destructive earthquakes occur along these faults, which are of strike-slip character (Ergin et al., 1967; Over et al., 2002). Both historical and instrumental records reveal that this region has been affected by devastating earthquakes for nearly 2000 years (Willis, 1928; Sieberg, 1932; Ergin et al., 1967; Ambraseys, 1970; Poirier and Taher, 1980; Soysal et al., 1981; Ambraseys and Barazangi, 1989). Important earthquakes that occurred in the historical period on the segments forming the EAFZ were respectively 29 November 1114 (*M* > 7.8), 28 March 1513 (*M* > 7.4), 1822 Antakya Earthquake ( $M<sub>s</sub> = 7.5$  approximately 200 km surface rupture), 1866 Karlıova-Bingöl Earthquake (M<sub>s</sub>)

= 7.2 approximately 45 km surface rupture), 1872 Amik Lake earthquake ( $M<sub>S</sub> = 7.2$  approximately 20 km surface rupture), 1874 and 1875 Hazar Lake earthquakes ( $M_{\rm s}$  = 7.1 and  $M<sub>S</sub> = 6.7$ , 45 and 20 km surface rupture) and 2 March 1893 Malatya Earthquake (M<sub>s</sub> = 7.1) (Ambraseys and Jackson, 1998) (please see details in supplementary material; Table S1). In the instrumental period 04.12.1905  $(M<sub>s</sub> = 6.8), 28.09.1908$  ( $M<sub>s</sub> = 6.1), 20.08.1966$  ( $M<sub>s</sub> = 6.2$ ),  $22.05.1971 \ (M<sub>s</sub> = 6.8), 27.01.2003 \ (M<sub>d</sub> = 6.1), 08.03.2010$  $(M<sub>L</sub> = 5.8)$ , and 24.01.2020 ( $M<sub>W</sub> = 6.8$ ) earthquakes did occur, but these earthquakes were not as destructive as historical earthquakes (please see details in supplementary material; Table S2). Çetin et al. (2003) interpreted the seismic quiescence along the entire EAFZ as that the fault was locked for this period. Historical earthquakes start from the NE end of the EAFZ and continue toward the SW. The general distribution of historical earthquakes is concentrated in the middle and SW parts of the EAFZ (Figure 3).



and https://deprem.afad.gov.tr/event-instrumental). Orange stars symbolize destructive earthquakes. The map also depicts **Figure 3.** Historical (between 69–1895 AD) and instrumental (between 01 January 1900 –05 February 2023) earthquakes in and around the Eastern Anatolian Fault Zone (compiled from AFAD website (https://deprem.afad.gov.tr/event-historical the fault segments of the East Anatolian Fault Zone (*1: Karlıova segment, 2: Ilıca segment, 3: Palu segment, 4: Pütürge segment, 5: Erkenek segment, 6: Pazarcık segment, 7: Amanos segment*).

On the Gölbaşı-Türkoğlu segment, where the fault folds into the SW, there has not been a devastating earthquake for the last 500 years to cause a surface rupture.

GPS studies show a slip rate of 9–10 mm/year for the EAFZ (McClusky et al., 2000; Reilinger et al., 2006; Aktug et al., 2016). About one-third of this is shared by the northern coast after bifurcation (Westaway, 2004; Altunel et al., 2009; Mahmoud et al., 2013; Emre et al., 2013).

The last devastating earthquake in this region occurred in 1872 and its acceleration was estimated to be 0.4 g (Över et al., 2004). After the previous earthquakes along the EAFZ and the  $M_w = 6.8$  earthquake in Sivrice (Elazığ) in 2020, it was stated that an earthquake was expected by pointing out a seismic gap in this area, on which many researchers agree (Nalbant et al., 2002; Duman and Emre, 2013; Alkan et al., 2021).

### **3. Stress transfer**

In fact, the phenomenon called stress transfer, which is considered under the name of triggering after an earthquake occurs at a point, is defined as the transfer of the energy of a moving part to another part in conditions where there is more than one fault system in a fault zone. The most common representation of elastic stress interaction is based on the calculation of Coulomb stress changes caused by earthquake dislocations (Harris, 1998; King and Cocco, 2001). Coulomb stress variations are widely used in the literature to investigate fault interactions between major earthquakes and also to model aftershock models and seismicity rate changes over longer time windows. Despite the large number of observations that trigger earthquakes, many questions remain unanswered about the physical mechanisms that could explain earthquake formation and the spatio-temporal pattern of seismicity. In particular, it is well known that other time-dependent processes such as aftershock, viscoelastic relaxation of the lower crust,

and fluid flow can alter the induced coseismic stress field. These stress perturbations occur at different temporal scales. By limiting the analysis to the aftershock duration, pore fluid flow can play a dominant role in inducing seismicity. Propagation processes of pore pressure relaxation in fractured and saturated rocks have been proposed to explain both earthquakes and triggered seismicity (Shapiro et al., 2003; Antonioli et al., 2005).

The focal mechanism solution of the earthquakes, the details of which are given in Table, showed that the earthquakes in Pazarcık and Elbistan segments belong to strikeslip, the segment where Yeşilyurt earthquake is located to dip-slip reverse faulting, Göksun earthquake to normal faulting, and Defne earthquake to strike-slip faulting (Figure 4). More than 9000 aftershocks have occurred from these earthquakes.

#### **4. Coulomb stress analysis**

Investigation of areas that are likely to experience earthquakes in the future is one of the important stages of earthquake hazard determination. To reach this goal, the processes that create the earthquake should be well known. Stress changes in active faults due to earthquakes are decisive for the estimation of dangerous zones. An earthquake can lead to the normal occurrence of earthquakes that may occur after it, delay it, or trigger other faults (Stein et al., 1997; Çakır et al., 2003; Nalbant, 2005). Stress accumulation occurs in fault zones depending on tectonic processes, and these stresses decrease with the occurrence of earthquakes. When there is an earthquake, the earthquakes that occur with the movement in the relevant fault reduce the tension in the region. After the stress reduction, the earthquake hazard decreases until a new stress accumulation occurs (Chinery, 1963). The increase in the tension in the region due to the earthquake action causes the nearby faults to be triggered. The decrease or accumulation of ten-

N <sub>o</sub>	Date (UTC)		Latitude (°N) Longitude (°E)   Depth (km)		Magnitude Type		Location	$S/D/R$ (°)	Source
$\mathbf{1}$	$06 - 02 - 2023$ 01:17:32	37.288	37.043	8.6	$M_{\rm w}$	7.7	Pazarcık (K.Maras)	233/74/18	AFAD
$\overline{2}$	$06 - 02 - 2023$ 01:28:16	37.304	36.920	6.2	$M_{\rm w}$	6.6	Nurdağı (Gaziantep)	$187/43/-30$	AFAD
$\overline{3}$	$06 - 02 - 2023$ 10:24:47	38.089	37.239	7.0	$M_{\rm w}$	7.6	Elbistan (K.Maras)	90/86/13	<b>AFAD</b>
$\overline{4}$	$06 - 02 - 2023$ 10:51:28	38.311	38.191	7.0	$M_L$	5.7	Yeşilyurt (Malatya)	241/33/107	<b>AFAD</b>
5	$06 - 02 - 2023$ 12:02:11	38.071	36.478	17.04	$M_{\rm w}$	5.9	Göksun (K.Maras)	$9/56/-95$	<b>AFAD</b>
6	20-02-2023 17:04:28	36.121	36.074	16.74	$M_{\rm W}$	6.4	Defne (Hatay)	$214/57/-44$	<b>AFAD</b>

**Table.** Source parameters of six induced earthquakes. The fault plane solutions are compiled from the AFAD website (https://deprem. afad.gov.tr/event-catalog).



**Figure 4.** Epicenter distribution of the Pazarcık earthquake ( $M_w$  = 7.7), Nurdağı earthquake ( $M_w$  = 6.6), Elbistan earthquake ( $M_w$ ∆ fault plane solutions of mainshocks are taken from the AFAD website (https://deprem.afad.gov.tr/event-catalog). The earthquake = 7.6), Yeşilyurt earthquake ( $M_L$  = 5.7), Göksun earthquake ( $M_w$  = 5.9), and Defne earthquake ( $M_w$  = 6.4), and their aftershocks from February 06, 2023 to February 27, 2023. The magnitudes of the earthquakes are depicted with different circular symbols. The magnitudes of the earthquakes are depicted with different circular symbols. The database is compiled from Kandilli Observatory and Earthquake Research Institute (http://www.koeri.boun.edu.tr/sismo/2/en/).

sion in the region is possible by observing the Coulomb coefficient in this study was assumed as  $\mu' = 0.4$ stress variation.

Earthquake and fault relationship can be established and the earthquake hazard in the region can be calculated by determining the Coulomb stress variation (Toda et al., 1998). Displacement in an active fault with the potential to produce earthquakes causes static stress changes. The displacement in this elastic half-space is used to calculate the the application of the AFAD website (https://deprement in this elastic half-space is used to calculate the  $\frac{2016}{6}$ . Stress triggering describes the sensiti 3D stress field and multiplied by the elastic stiffness to derive the stress changes. The Coulomb stress variation de-<br>by extrinsic deformation events (King et al., 19 pends on the geometry and slippage of the earthquake, the geometry of the fault and the sense of slip, and the effective friction coefficient (Stein et al., 1994). The Coulomb failure stress change (**ΔσCFS**) caused by a source earthquake on an adjacent receiver fault can be expressed as

$$
\Delta \sigma_{CFS} = \Delta \tau + \mu' \Delta \sigma_n \tag{1}
$$

 $\mu'$  is the effective friction coefficient. The effective friction where  $\Delta \sigma$ <sub>CFS</sub> is the critical Coulomb failure stress on a receiving fault, is the shear stress change on a given fault plane (assumed positive in the direction of fault slip), is the normal stress change (positive when unclamped) and ∆ () coefficient is dimensionless and varies between 0.2 and 0.8 for the continental structures. The effective friction

coefficient in this study was assumed as  $\mu' = 0.4$  due to etermining the Coulomb stress variation (10da et al.,  $\qquad$  1994; Lin and Stein, 2004; Toda et al., 2011). When a fault faults resulting from increases in Coulomb stress caused the continental strike-slip fault mechanisms. In addition, arthquake and fault relationship can be established Young's modulus was used as  $\vec{E} = 80$  GPA with Poisson's ronment is deformed, and the stress field changes (Ansari, ratio of  $m = 0.25$  for the elastic half-space (King et al., slip occurs during an earthquake, the surrounding envi-2016). Stress triggering describes the sensitive rupture of by extrinsic deformation events (King et al., 1994). Although adjacent displacements generally yield small-scale stress changes, the fields of deteriorated Coulomb stress states have been successfully used to explain the distribution of stress-induced aftershocks. On June 28, 1992, approximately three h after an earthquake (*M* = 7.3) near Landers, California, 40 km away, the Big Bear aftershock (*M* = 6.5) occurred. The calculated Coulomb stress changes from both of these earthquakes showed a western lobe increasing Coulomb stress of 2.1–2.9 bar, which was due to the displacement associated with both earthquakes. More than 75% of the approximately 20,000 aftershocks within a 5 km radius after 25 days occurred in areas of increased Coulomb tension and the remainder in areas of decreased Coulomb tension. Another example, though extended over a longer period of time, was 13 earthquakes (*M* ≥ 6.6)

along the North Anatolian Fault System from 1939 to 1999, which occurred along Turkey's NAFZ.

Eleven of these earthquakes occurred in areas of increased Coulomb tension caused by a previous rupture (Stein et al., 1997; Barka et al., 1999). Similarly, it is a known fact that earthquakes occurring along the East Anatolian Fault Zone are segments that trigger each other and create stress.

#### **5. Results and discussions**

The transfer of the energy of a moving segment to another segment after an earthquake occurs in one segment in a multi-segment fault system has become a common phenomenon observed in many parts of the world. NAFZ and EAFZ are two well-known examples of these situations in Turkey. The direction and orientation of elastic stress interaction can be determined by calculating Coulomb stress changes. The results of the Coulomb stress analysis performed after the 24 January 2020 Sivrice-Elaz**ığ** earthquake  $(M<sub>w</sub> = 6.8)$  clearly showed that the stress was transferred to the next segments, Adıyaman and Kahramanmaraş in the southwest (Alkan et al., 2021). The occurrence of very destructive earthquakes in these segments after about three years is understood as an expected situation in this sense. Stress modeling of six earthquakes, which started with the earthquake  $(M_v = 7.7)$  in Pazarcık-Kahramanmaraş on 06 February 2023 and triggered each other for very short periods, was carried out with Coulomb 3.3 software (Toda et al., 2011). The important parameters to calculate the Coulomb stress change of an earthquake are the location (latitude and longitude in degree), depth (km), magnitude, dip (degree), strike (degree), and rake (degree). These parameters are very vital to eliminate uncertainties regarding the parameters of the Coulomb stress analysis. For three-component seismograms (Z, N, and E) of a large earthquake, it is generally a less complex problem to mark the P- and Sphases, determine the maximum amplitudes of the P- and S-phases, and specify the coda length of the seismograms. Therefore, the focal mechanism solutions calculated by the AFAD or other seismology institutes such as the USGS or the KOERI can be reliable. In this study, hence, we created the Coulomb stress change maps and cross-sections using the focal mechanism solutions of AFAD.

The results show that the earthquakes after the first main shock occurred in the areas where Coulomb stress increased. Coulomb stress analysis performed at four different depths (5, 10, 15, and 20 km) for the first main shock  $(M_w = 7.7)$  and the aftershock  $(M_w = 6.6)$  that occurred afterward, indicates that the stress was increased Malatya-Adıyaman in the northeast, Elbistan-Göksun in the northwest, Hatay-Syria border in the south and İskenderun Bay in the southwest directions (Figure 5). Especially in the shallow depths of the crust  $(-10 \text{ km})$ , it can be obser-

ved that the stress values are quite high  $(\sim 0.1$  bar) both along the fault zone and along the four different directions mentioned. On the contrary, the stress values along the city centers of Adana and Osmaniye, east of Gaziantep and north of Kahramanmaraş are low (~-0.1 bar) between 0–20 km depth. On the other hand, positive stresses along the Yumurtalık fault and the Karataş fault were observed to be transferred towards the Cyprus trench in the southwest. After the first two earthquakes ( $M_w = 7.7$  and  $M_w = 6.6$ ), the main shock ( $M_w$  = 7.6) that occurred and was triggered in Ekinözü-Elbistan (Kahramanmaraş) and the other earthquakes Yeşilyurt (Malatya) ( $M_{\rm L}$  = 5.7), Göksun (Kahramanmaraş)  $(M_w = 5.9)$  (Figure 6) and Defne-Samandağ (Hatay)  $(M_v = 6.4)$  earthquakes occurred in regions where stress is positive (Figure 7). Coulomb stress analysis of the Ekinözü-Elbistan ( $M_{\rm w}$  = 7.6), Yeşilyurt-Malatya ( $M_{\rm L}$  = 5.7), and Göksun-Kahramanmaraş ( $M_w$  = 5.9) earthquakes, which are the next three earthquakes occurring in different segments following the  $M_w = 7.7$  main shock and the  $M_w = 6.6$  aftershock, was performed and it was clearly seen that it was transferred to the **Çüngüş**, Palu, and Pazarcık Segments located between Elazığ and Diyarbakır in the east and to the Sarız Fault and Ecemiş Fault in the west. The aftershock sequences that occur after the main shocks also confirm this energy transfer. On the other hand, as an important note, a stress increase was observed towards Malatya and Ovacık faults at 0–20 km depth in the north and between Gaziantep and Kilis in the south (Figure 6).

The current stress situation has become clear as a result of the combined Coulomb analysis of six earthquakes, which were included in the Defne-Hatay earthquake  $(M_w = 6.4)$  that occurred on the 20th of February, 2023. According to this, the stresses  $(-0.1 \text{ bar})$  were positively transferred regions observed at a depth of 0–20 km around the İskenderun Bay in the southwest, towards Karlıova triple junction along Elazığ-Malatya in the northeast, along the Kahramanmaraş fault zone between Kahramanmaraş-Göksun in the west, along the Hatay-Syria border in the south, that is, in the region where the segments of the EAFZ and DSFZ intertwine. Considering the epicenters of more than 9000 aftershocks, it can be assumed that stress transfers will continue in these directions (Figure 7).

In addition, the vertical displacements of the focal mechanism solution of  $M_w = 7.7$  and  $M_w = 7.6$  magnitudes of the two mainshock solutions using Coulomb 3.3 software (Toda et al., 2011) are given in Figure 8 and Figure 9, respectively, and the horizontal displacement diagrams are given in Figure 10a and 10b, respectively. When Figure 8 and Figure 9 are examined, vertical crosssections were obtained for the 0–30 km depth range along the A-B profiles. In the sections formed along the fault zones, the yellow circles represent the epicenter point in both earthquakes.



Figure 5. Coulomb stress changes due to the first main shock and aftershock ( $M_{\text{w}}$ ,  $M_{\text{max}}$  and  $M_{\text{w}}$  or  $\sigma$ ) seedfied after and after pairs of the depths of 5.0, 10.0, 15.0, and 20.0 km. The big yellow circles locations of aftershocks that occurred from 06.02.2023 at 01:17:32 (UTC) to 06.02.2023 at 10:24:47 (UTC). Claret red circles show<br>https://www.characteristics.com/2020/2023/at 01:17:32 (UTC) to 06.02.2023 at 10:24:47 (UTC). **Figure 5.** Coulomb stress changes due to the first main shock and aftershock ( $M_w$  = 7.7 and  $M_w$  = 6.6) occurred around the mechanism results taken from the AFAD website (https://deprem.afad.gov.tr/event-catalog). The orange circles depict epicentral the city centers. Black lines represent the active faults (modified from Emre et al. (2018)).

The first main shock coincides with the Narlı Segment According to the solutions obtained, the characteristic (https://deprem.afabrications.gov.tr/event-catalog). The AFAD website (https://deprem.afabrications.gov.tr/event at the northern end of the DSFZ, while the second main shock coincides with the Çardak Fault of the EAFZ. Considering the general stress characteristics in both stress directions, SW-NE for the first main sections, low stress around the hypocentral region and high-stress transfer along the fault zone in the NE- main shock occurred. It was observed that SW direction for the first main shock and in the E-W direction for the second main shock emerged. The stress zones where the characteristic of these two main shocks generally reflects the entire energy system of the region. On the other hand, horizontal displacement vectors obtained from focal mechanism solutions of the Disaster and Emergency Management Authority (AFAD) (Table) for both main shocks are shown in Figures 10a and 10b. Besides, in the horizontal displacement calculation, the medium, which controls the environmental propagation of the earthquake source's impact, is assumed uniform (Yildiz et al., 2021).

According to the solutions obtained, the character of both earthquakes supports the solution of the left-sided focal mechanism. In addition, similar to positive Coulomb stress directions, SW-NE for the first main shock and E-W directional horizontal displacements for the second main shock occurred. It was observed that horizontal displacements reached  $~6$  m, especially along the fault zones where the rupture occurred (green continuous line).

#### **6. Conclusions**

Modeling results based on Coulomb stress analysis after the earthquakes defined as 06 February 2023 Kahramanmaraş earthquakes explain the formation of the earthquake sequence occurring in the EAFZ of the elastic stress transfer. Especially the fact that the earthquake occurred at a shallow depth caused higher energy to reach the surface and the stress transfer to take place in a short time. How-



**Figure 6.** Coulomb stress changes due to the earthquakes ( $M_w = 7.6$ ,  $M_L = 5.7$ , and  $M_w = 5.9$ ) that occurred around the  $M_w$  =  $M_w = 5.9$  for the explorer of the  $M_w$  =  $M_w = 5.9$ orange circles depict epicentral locations of aftershocks that occurred from 06.02.2023 at 01:17:32 (UTC) to 06.02.2023 at Kahramanmaraş and Malatya for the depths of 5.0, 10.0, 15.0, and 20.0 km. The big yellow circles show the epicenter of the earthquakes with the focal mechanism results taken from the AFAD website (https://deprem.afad.gov.tr/event-catalog). The 10:51:30 (UTC). Claret red circles show the city centers. Black lines represent the active faults (modified from Emre et al. (2018)).

ever, the only and most important reason for the very high damage caused by earthquakes was not the complex stress between Elazi<mark>g-Diyarbakir in the northea</mark> faulting system, but mostly old and unstable construction which was not built in accordance with the soil conditions.  $\qquad \quad$  south. Then, on February 20, 2023, an earthqua

In order to follow the change in stress during the study and to understand whether the stress failure is as expected ated as the failure of the tension in the direction in a very short time, first of all, Coulomb stress analysis of the earthquakes ( $M_v$  = 7.7) and ( $M_w$  = 6.6) was made for 5, 10, 15, and 20 km depths and was observed that the stress distribution at the time  $\epsilon$ stresses increased in the direction of Malatya-Adıyaman in the northeast, Elbistan-Göksun in the northwest, Hatay-Syria border in the south and **İ**skenderun Bay in the southwest. Coulomb analysis was made for 5, 10, 15, and 20 km depths for earthquakes that occurred in Ekinözü-Elbistan (Kahramanmaraş) ( $M_w$  = 7.6) in the west, Yeşilyurt (Malatya)  $(M<sub>L</sub> = 5.7)$  in the direction of Malatya-Adıyaman and Göksun (Kahramanmaraş)  $(M_{w} = 5.9)$  in the west direction of Göksun-Elbistan where the stresses in the pre-

vious analysis increased and were observed increased in stress between Elazığ-Diyarbakır in the northeast, between Malatya-Elazığ in the north and Gaziantep-Kilis in the south. Then, on February 20, 2023, an earthquake of  $M_w$  = 6.4 in Defne-Hatay occurred and this earthquake was evaluated as the failure of the tension in the direction of the Hatay-Syria border in the south. After that, a combined Coulomb analysis of six earthquakes is made and the current stress distribution at the time of this study is observed as the regions where stresses are transferred between Elaz**ığ** and Malatya in the northeast, Kahramanmaraş-Göksun in the west, Hatay and Syria border in the south, and **İ**skenderun Bay in the southwest. The vertical and horizontal displacements of the focal mechanism solution of two main shocks with  $M_w = 7.7$  and  $M_w = 7.6$  magnitudes also provided confirmatory results in terms of showing the variation of stress directions and their depths.

The absence of a major earthquake in this segment of



circles show the epicenter location of the earthquakes with the focal mechanism results taken from the AFAD website depict epicentral locations of aftershocks that occurred from February 06, 2023, to February 23, 2023. Claret red circles **Figure 7.** Coulomb stress changes of all induced shocks for the depths of 5.0, 10.0, 15.0, and 20.0 km. The big yellow (https://deprem.afad.gov.tr/event-catalog). The number above each earthquake is given in Table. The orange circles show the city centers. Black lines represent the active faults (modified from Emre et al. (2018)).



Figure 8. The vertical cross-section of Coulomb stress changed along profile A-B is taken from the AFAD website (https://deprem.afad.gov.tr/event-catalog). during the Pazarcık earthquake ( $M_v$  = 7.7) at a depth of 8.6 km. The catalog information



depth of 7.0 km. The catalog information is taken from the AFAD website (https://deprem.afad.gov.tr/event-catalog). **Figure 9.** The vertical cross-section of Coulomb stress changed along profile A-B during the Elbistan earthquake ( $M_w$  = 7.6) at a



 $(M_{_{\rm w}}$  = 7.6). Blue lines depict active fault zones taken by Emre et al. (2018). **Figure 10.** The horizontal displacement vector analyses for (a) the Pazarcık earthquake ( $M_w$  = 7.7) and (b) the Elbistan earthquake

the EAFZ for a long time has caused great tension in this region and the rupture to occur with very high energy.

#### **Acknowledgments**

We thank the anonymous reviewer for his/her interest in this work and for critically reviewing the manuscript. The authors thank the Kandilli Observatory and Earthquake

#### **References**

- AFAD (2023). The Disaster and Emergency Management Authority (AFAD) Earthquake Catalog. AFAD | deprem.gov.tr. Access date: 20.02.2023
- Aktug BA, Ozener H, Dogru A, Sabuncu A, Turgut B et al. (2016). Slip rates and seismic potential on the East Anatolian Fault System using an improved GPS velocity field. Journal of Geodynamics 1(94): 1-2. https://doi.org/10.1016/j.jog.2016.01.001
- Alkan H, Büyüksaraç A, Bektaş Ö, Işık E (2021). Coulomb stress change before and after 24.01. 2020 Sivrice (Elazığ) Earthquake (Mw = 6.8) on the East Anatolian Fault Zone. Arabian Journal of Geosciences 14 (23): 1-12. https://doi.org/10.1007/s12517- 021-09080-1
- Altunel E, Meghraoui M, Karabacak V, Akyüz SH, Ferry M et al. (2009). Archaeological sites (tell and road) offset by the Dead Sea Fault in the Amik Basin, southern Turkey. Geophysical Journal International 179 (3): 1313-1329. https://doi. org/10.1111/j.1365-246X.2009.04388.x
- Ambraseys NN (1970). Some characteristic features of the Anatolian fault zone. Tectonophysics 9 (2-3): 143-165.
- Ambraseys NN, Barazangi M (1989). The 1759 earthquake in the Bekaa Valley: implications for earthquake hazard assessment in the Eastern Mediterranean region. Journal of Geophysical Research Solid Earth B4: 4007-4013. https://doi.org/10.1029/ JB094iB04p04007.
- Ambraseys NN, Jackson JA (1998). Faulting associated with historical and recent earthquakes in the Eastern Mediterranean region. Geophysical Journal International 133: 390-406.
- Ansari S (2016). Co-seismic stress transfer and magnitudefrequency distribution due to the 2012 Varzaqan-Ahar earthquake doublets (Mw 6.5 and 6.4), NW Iran. Journal of Asian Earth Sciences 132: 129-137. https://doi.org/10.1016/j. jseaes.2016.10.006
- Antonioli A, Piccinini D, Chiaraluce L, Cocco M (2005). Fluid flow and seismicity pattern: evidence from the 1997 Umbria-Marche (central Italy) seismic sequence. Geophys Research Letters 32: L10311. https://doi.org/10.1029/2004GL022256
- Barka AA (1999). The 17 August 1999 Izmit earthquake. Science 285: 1858-1859.
- Bayrak E, Bayrak Y, Yilmaz Ş, Softa M, Turker T (2015). Earthquake hazard analysis for East Anatolian Fault Zone, Turkey. Natural Hazards 76: 1063-1077. https://doi.org/10.1007/s11069-014- 1541-5

Research Institute (KOERI) and the Disaster and Emergency Management Authority (AFAD) for providing earthquake catalogs. Some of the figures are plotted by using the Generic Mapping Tools (Wessel et al., 2019) software package. The active fault database is taken from Emre et al. (2018).

- Bozkurt E (2001). Neotectonics of Turkey-a synthesis. Geodinamica Acta 14 (1-3): 3–30.
- Bulut F, Bohnhoff M, Eken T, Janssen C, Kılıç T et al. (2012). The East Anatolian Fault Zone: Seismotectonic setting and spatiotemporal characteristics of seismicity based on precise earthquake locations. Journal of Geophysical Research 117: B07304. https://doi.org/10.1029/2011JB008966
- Chinnery MA (1963). The stress changes that accompany strike-slip fault- ing. Bulletin of the Seismological Society of America 53: 921-932.
- Çetin H, Güneyli H, Mayer L (2003). Paleoseismology of the Palu-Lake Hazar Segment of the East Anatolian Fault Zone, Turkey. Tectonophysics 374: 163-197. https://doi.org/10.1016/j. tecto.2003.08.003
- Çakır Z, Barka AA, Evren E (2003). Coulomb Stress Interactions and the 1999 Marmara Earthquakes. Turkish Journal of Earth Sciences 12: 1. https://journals.tubitak.gov.tr/earth/vol12/ iss1/6
- Duman TY, Emre Ö (2013). The East Anatolian Fault: geometry, segmentation and jog characteristics. Geological Society London Special Publications 372 (1): 495-529. https://doi. org/10.1144/SP372.14
- Emre Ö, Duman TY, Özalp S, Şaroğlu F, Olgun Ş et al. (2018). Active fault database of Turkey. Bulletin of Earthquake Engineering 16 (8): 3229-3275. https://doi.org/10.1007/s10518-016-0041-2
- Ergin K, Guglu U, Uz Z (1967). A catalogue of earthquakes of Turkey and surrounding area (11 A.D. to 1964 A.D.). Maden Fakültesi Arz Fiziği Enstitüsü Yayin Istanbul 24: 169.
- Gülerce Z, Shah ST, Menekşe A, Özacar AA, Kaymakci N et al. (2017). Probabilistic Seismic‐Hazard Assessment for East Anatolian Fault Zone Using Planar Fault Source Models. *Bulletin of the Seismological Society of America* 107 (5): 2353-2366. https:// doi.org/10.1785/0120170009
- Harris R (1998). Introduction to special section: Stress triggers, stress shadows, and implications for seismic hazard. Journal of geophysical research 103 (10): 24347-24358.
- Irmak TS, Toker M, Yavuz E, Şentürk E, Güvenaltın MA (2021). New insight into the 24 January 2020, Mw 6.8 Elazığ earthquake (Turkey): An evidence for rupture-parallel pull-apart basin activation along the East Anatolian Fault Zone constrained by Geodetic and Seismological data. Annals of Geophysics 64 (4): SE439-. https://doi.org/10.4401/ag-8638
- Kartal RF, Kadirioğlu FT (2013). Doğu Anadolu Fayının sismotektoniği ve bu fay üzerindeki son beş yıllık deprem aktivitesinin istatistiksel analizi. 66. Türkiye Jeoloji Kurultayı, 01-05 Nisan 2013, ODTÜ Kültür ve Kongre Merkezi, Ankara (in Turkish).
- King GC, Stein RS, Lin J (1994). Static stress changes and the triggering of earthquakes. Bulletin of the Seismological Society of America 84 (3): 935-953. https://doi.org/10.1785/BSSA0840030935
- King GCP, Cocco M (2001). Fault interaction by elastic stress changes: new clues from earthquake sequences. Advances in Geophysics 44: 1-38. https://doi.org/10.1016/S0065-2687(00)80006-0
- Lin J, Stein RS (2004). Stress triggering in thrust and subduction earthquakes, and stress interaction between the southern San Andreas and nearby thrust and strike-slip faults. Journal of Geophysical Research B02303. https://doi. org/10.1029/2003JB002607
- Mahmoud Y, Masson F, Meghraoui M, Cakir Z, Alchalbi A et al. (2013). Kinematic study at the junction of the East Anatolian fault and the Dead Sea fault from GPS measurements. Journal of Geodynamics 67: 30-39. https://doi.org/10.1016/j.jog.2012.05.006
- McClusky S, Balassanian S, Barka A, Demir C, Ergintav S et al. (2000). Global Positioning System constraints on plate kinematics and dynamics in the eastern mediterranean and Caucasus. Journal of Gephysical Research 105: 5695- 5719.
- Nalbant SS, McCloskey J, Steacy S, Barka AA (2002). Stress accumulation and increased seismic risk in eastern Turkey. Earth and Planetary Science Letters 195 (3-4): 291-8. https:// doi.org/10.1016/S0012-821X(01)00592-1
- Nalbant S, Steacy S, Sieh K, Natawidjaja D, McCloskey J (2005). Earthquake risk on the Sunda trench. Nature 435: 756–757. https://doi.org/10.1038/nature435756a
- Okay AI, Tüysüz O (1999). Tethyan sutures of northern Turkey. Geological Society London Special Publications 156 (1): 475-515. https://doi.org/10.1144/GSL.SP.1999.156.01.22
- Over S, Ozden S, Yilmaz H (2004). Late Cenozoic stress state distributions along the Karasu Valley, SE Turkey. Tectonophysics 380: 43-68.
- Över S, Ünlügenç UC, Bellier O (2002). Quaternary stress regime change in the Hatay region (SE Turkey). Geophysical Journal International 148: 649-662.
- Öztürk S (2020). A study on the variations of recent seismicity in and around the Central Anatolian region of Turkey. Physics of the Earth and Planetary Interiors 30: 106453. https://doi. org/10.1016/j.pepi.2020.106453
- Poirier JP, Taher MA (1980). Historical seismicity in the near and Middle East, North Africa, and Spain from Arabic documents (VIIth-XVIIIth century). Bulletin of the Seismological Society of America 70 (6): 2185-2201.
- Reilinger R, McClusky S, Vernant P, Lawrence S, Ergintav S et al. (2006). GPS constraints on continental deformation in the Africa-Arabia-Eurasia continental collision zone and implications for the dynamics of plate interactions. Journal of Geophysical Research: Solid Earth 111 (B5): B05411. https:// doi.org/10.1029/2005JB004051
- Reilinger R, McClusky S, Vernant P, Lawrence S, Ergintav S et al. (2006). GPS constraints on continental deformation in the Africa-Arabia‐Eurasia continental collision zone and implications for the dynamics of plate interactions. Journal of Geophysical Research: Solid Earth 111 (B5).
- Reilinger R, McClusky S (2011). Nubia-Arabia-Eurasia plate motions and the dynamics of Mediterranean and Middle East tectonics. Geophysical Journal International 186 (3): 971-979. https://doi.org/10.1111/j.1365-246X.2011.05133.x
- Shapiro SA, Patzig R, Rothert E, Rindshwentner J (2003). Triggering of seismicity by pore pressure perturbations: permeability-related signature of the phenomenon. Pure and Applied Geophysics 160: 1051-1066. https://doi.org/10.1007/978-3-0348-8083-1\_16
- Sieberg A (1932). Erdbebengeographie. Handbuch der Geophysik 4: 708-744.
- Soysal H, Sipahioglu S, Kolcak D, Altinok Y (1981). Türkiye ve çevresinin tarihsel deprem katalogu. TUBITAK Proje no. TBAG 341 Istanbul 86 (in Turkish).Stein RS, King GC, Lin J (1994). Stress triggering of the 1994  $M = 6.7$  Northridge, California, earthquake by its predecessors. Science 265 (5177): 1432-1435.
- Stein RS, Barka AA, Dieterich JH (1997). Progressive failure on the North Anatolian fault since 1939 by earthquake stress triggering. Geophysical Journal International 128 (3): 594604. https://doi. org/10.1111/j.1365-246X.1997.tb05321.x
- Şengör AMC, Görür N, Şaroğlu F (1985). Strike-slip faulting and related basin formation in zones of tectonic escape: Turkey as a case study. in: Biddle K.T., ChristieBlick N. (Eds.), Strike-slip Faulting and Basin Formation, Society of Economic Paleontologists and Mineralogists 37: 227-264.
- Toda S, Stein RS, Reasenberg PA, Dieterich JH, Yoshida A (1998). Stress transferred by the 1995 Mw = 6.9 Kobe, Japan, shock: Effect on aftershocks and future earthquake probabilities. Journal of Geophysical Research Solid Earth 103 (B10): 24543-24565. https://doi.org/10.1029/98JB00765
- Toda S, Stein RS, Lin J (2011). Widespread seismicity excitation throughout central Japan following the 2011 M = 9.0 Tohoku earthquake and its interpretation by Coulomb stress transfer. Geophysical Research Letters 38 (7). https://doi. org/10.1029/2011GL047834
- Wessel P, Luis J, Uieda L, Scharroo R, Wobbe F et al. (2019). The Generic Mapping Tools Version 6. Geochemistry Geophysics Geosystems 20.
- Westaway R (2004). Kinematic consistency between the Dead Sea Fault Zone and the Neogene and Quaternary left-lateral faulting in SE Turkey. Tectonophysics 391: 203-237. https://doi.org/10.1016/j. tecto.2004.07.014
- Willis B (1928). Earthquakes in the Holy Land. Bulletin of the Seismological Society of America 18 (2): 73-103. https://doi. org/10.1785/BSSA0180020073
- Yildiz H, Cirmik A, Pamukcu O, Özdağ ÖC, Gönenç T et al. (2021). 12th June 2017 offshore Karaburun-Lesvos Island earthquake coseismic deformation analysis using continuous GPS and seismological data. Turkish Journal of Earth Sciences 30 (3): 341-358. https://doi.org/10.3906/yer-2008-3

## ALKAN et al. / Turkish J Earth Sci

#### **Supplementary Material for**

#### **Investigation of Earthquake Sequence and Stress Transfer in the Eastern Anatolia Fault Zone by Coulomb Stress Analysis**

Hamdi ALKAN1,\*, Aydın BÜYÜKSARAÇ2 ,Özcan BEKTAŞ3

## **Contents of this file**

**Tables S1 to S2:** 

 $\geq$  S1: Destructive earthquakes around the EAFZ in the historical period<br> $\geq$  S2: Earthquakes with a magnitude greater than 5.0 around the EAFZ is

 S2: Earthquakes with a magnitude greater than 5.0 around the EAFZ in the instrumental period

**Table S1.** Destructive earthquakes around the EAFZ in the historical period (compiled from AFAD and KOERI websites, respectively, https://deprem.afad.gov.tr/event-historical and http://www.koeri.boun.edu.tr/sismo/2/en/)





**Table S2.** Earthquakes with a magnitude greater than 5.0 around the EAFZ in the instrumental period (compiled from the USGS website https://www.usgs.gov/programs/earthquake-hazards/earthquakes)













